

Concepts and constraints for scaling from baseline metabolism to ecosystem carbon relations

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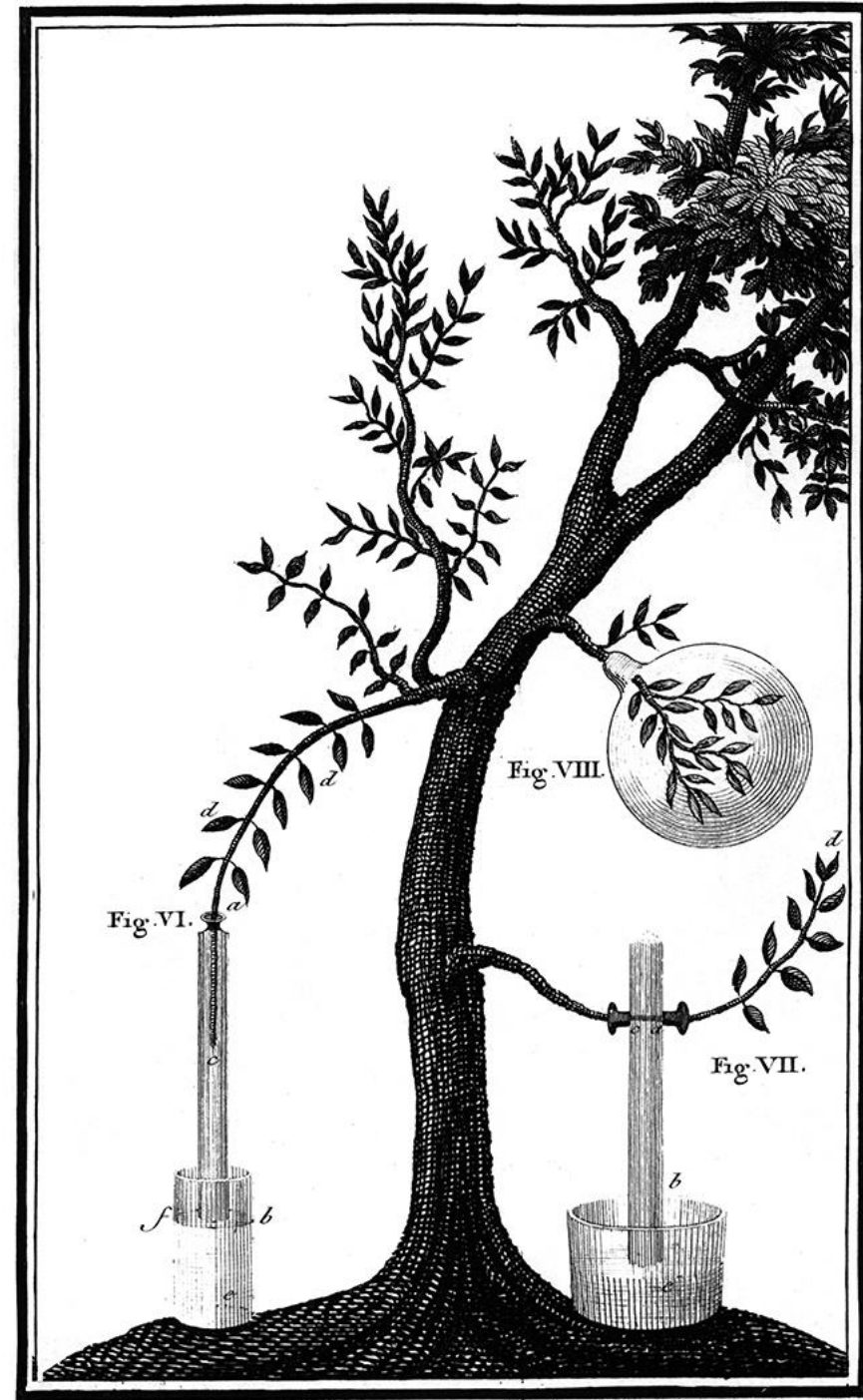


CLIMMANI/INTERFACE Workshop, Keflavik, Island June 15-17, 2011

In Geneva 200 years ago:

'The primary plant food comes from air'

Nicolas-Théodore de Saussure (1804)
Recherches chimiques sur la Végétation.
Paris.

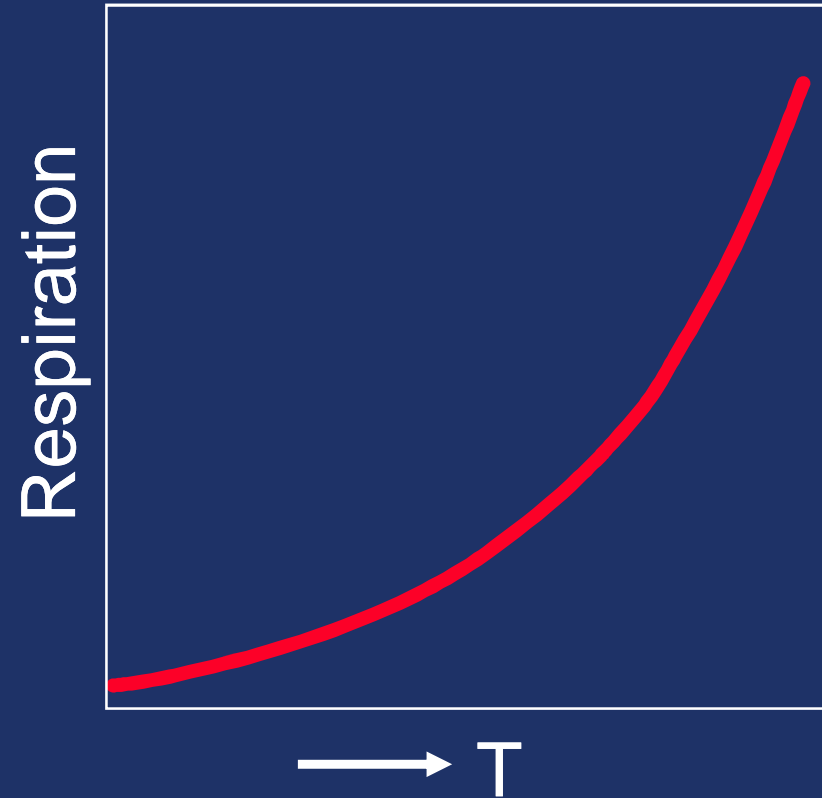
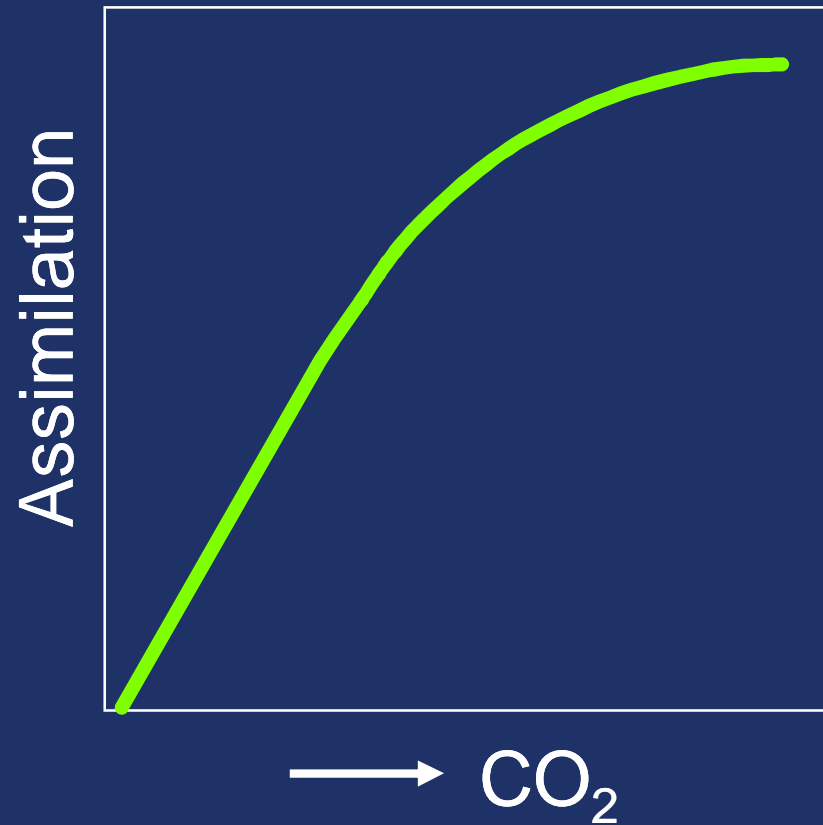


Gravé par Tardieu Laine, Rue de Sorbonne N. 383.

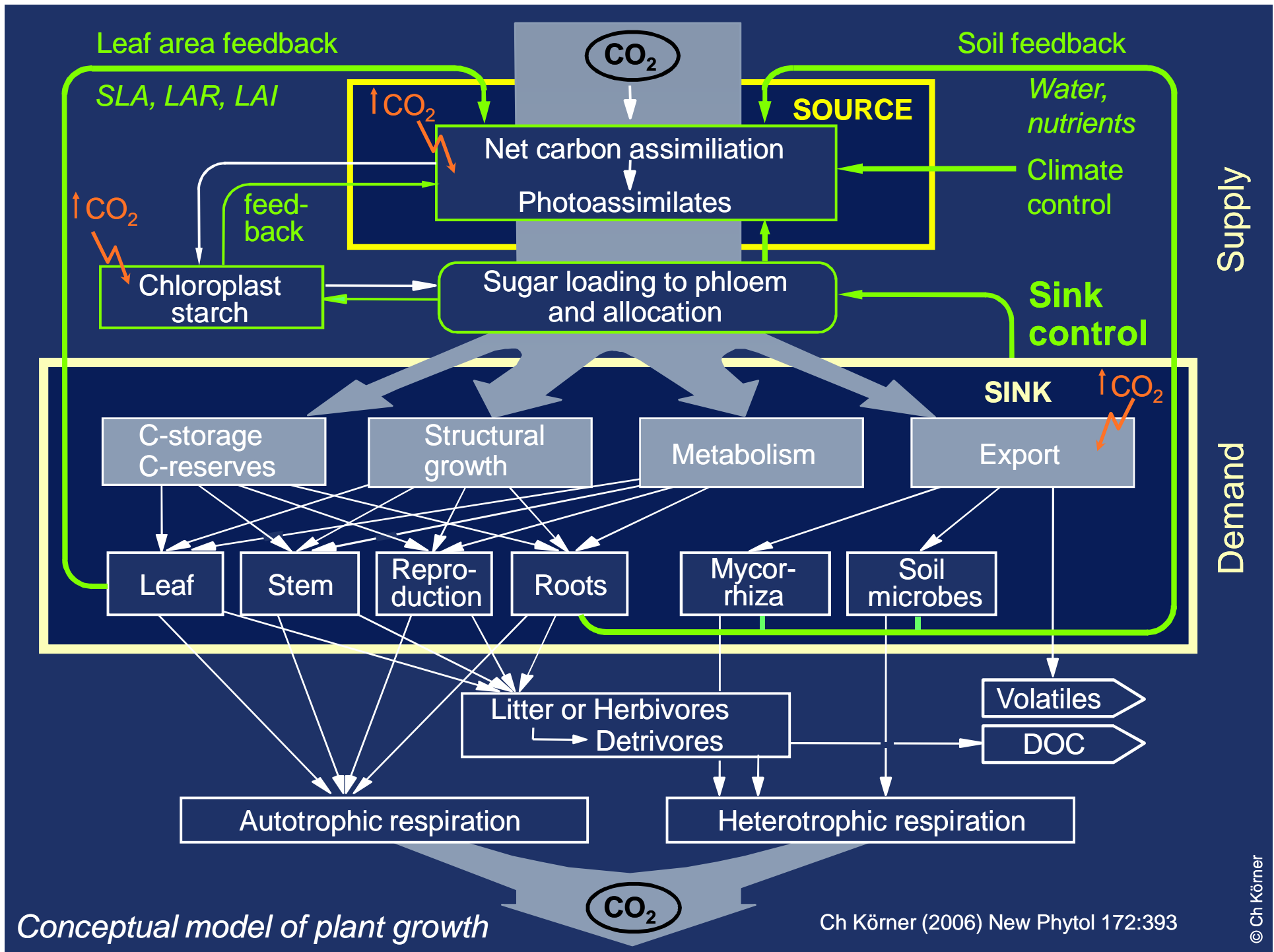
The new way.....



Are these basic physiological functions drivers or driven?



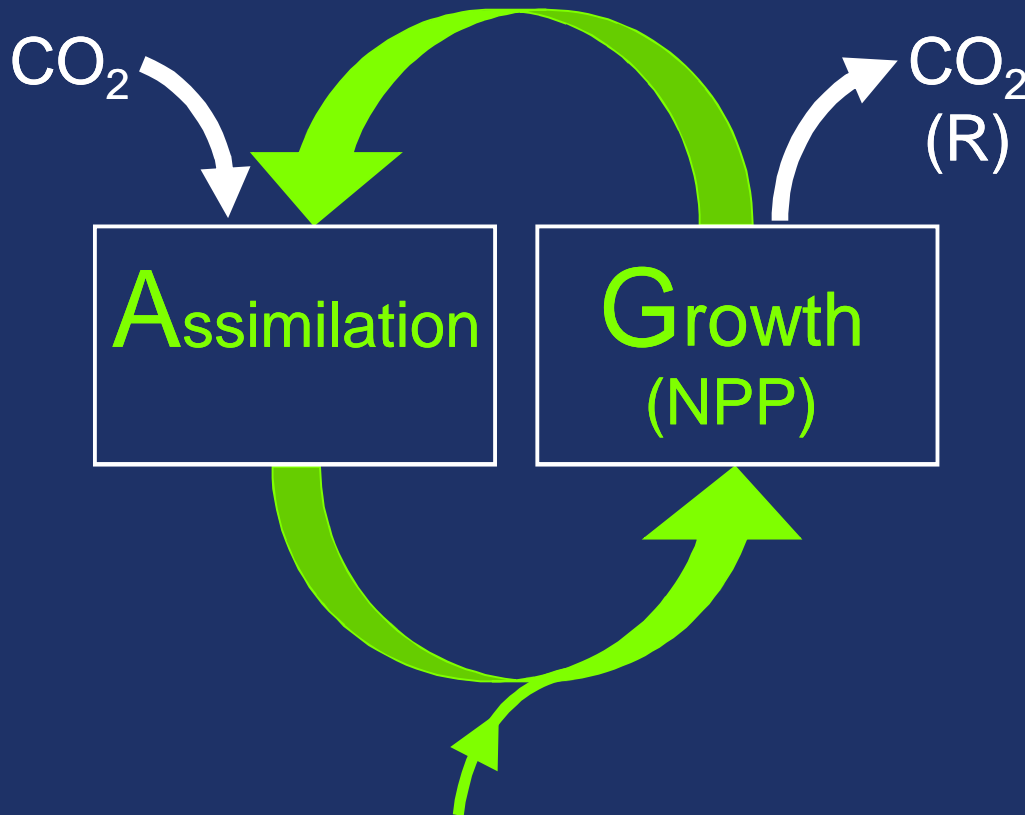
driven by what?



Assimilation drives growth or vice versa?

Light and CO₂ co-control source activity

Output is controlled by carbon input



Minerals etc.

(N input, water)

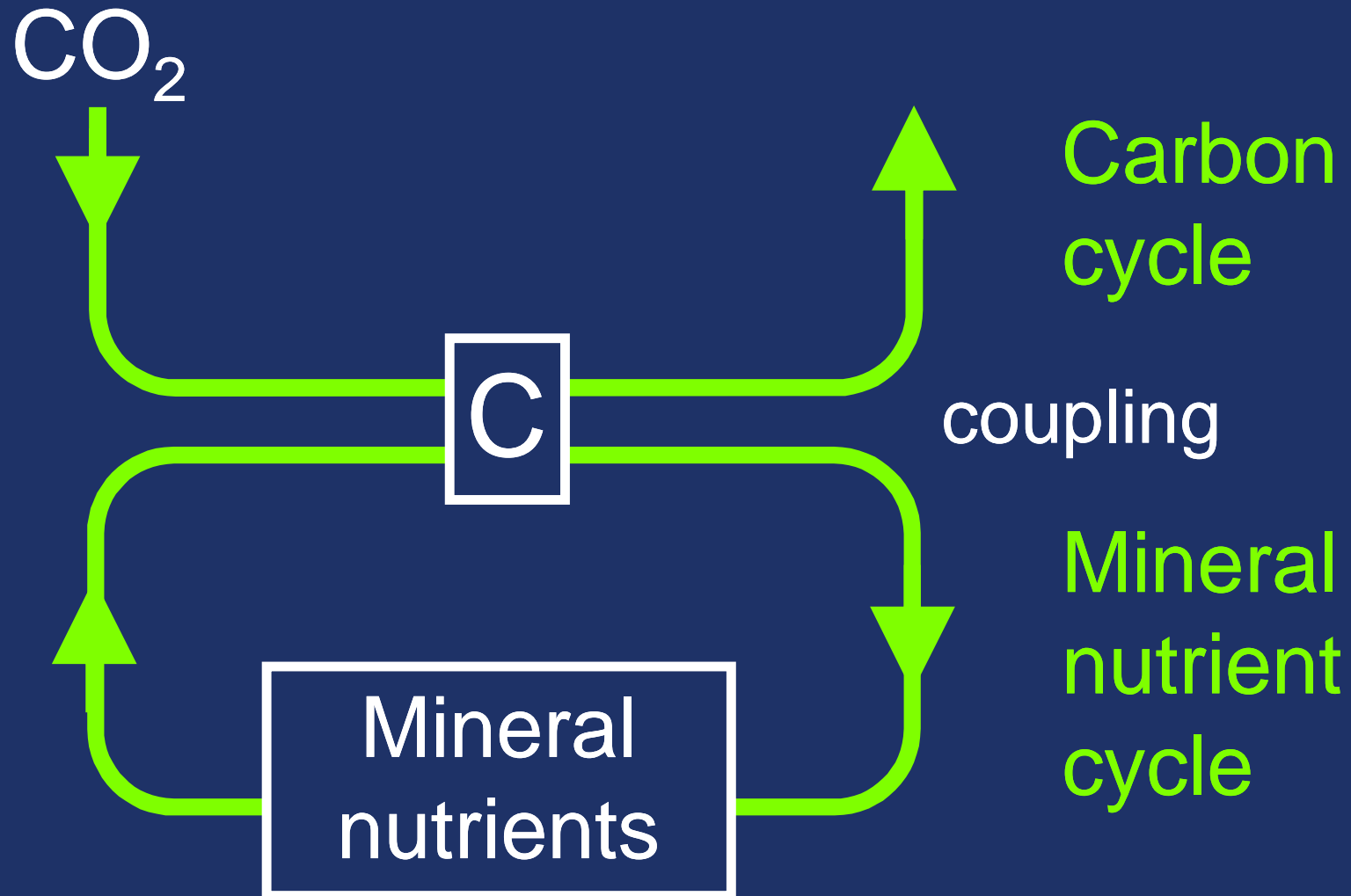
plus ontogeny & temperature
co-control sink activity


A or G first?

- Whenever M is limiting, G controls A
- Whenever CO₂ or light are limiting, A controls G
- The A ↔ G fine tuning is masking cause and consequence
→ manipulative experiments

chicken or egg first?

The ecosystem carbon cycle is driven by the nutrient cycle and other growth 'facilitators'





Experiments that lack a natural plant - rhizosphere - soil coupling are fundamentally unsuitable for making ecosystem scale inferences.

Annual ecosystem R is driven by

- demand (autotrophic R; growth, maintenance, nutrient uptake)
- substrate supply (heterotrophic R)
- duration of active period (environmental conditions x time)

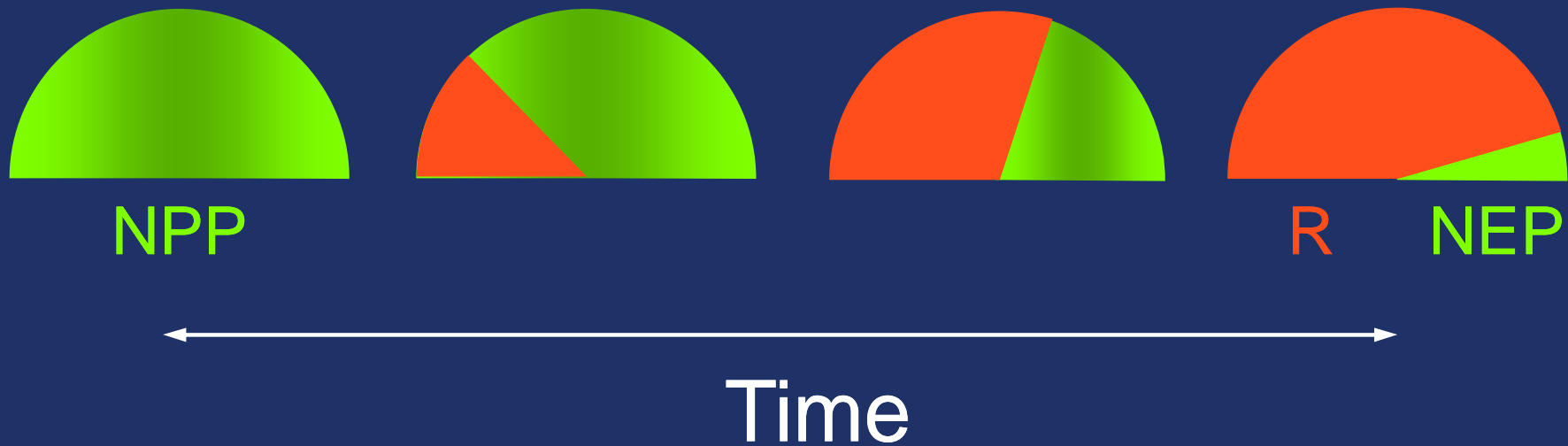
In the long run ...

- R is largely a function of NPP
- $NPP = f(\text{soils, age, temperature, etc.})$
- R cannot be decoupled from NPP
 - Hence, there is no independent temperature response.



you can eat
the lunch only once

Temporal shifts of $R_f(T, H_2O)$ should not be taken as evidence for increased or decreased ΣR



... but you may be eating it faster or slower

The consequences of climatic warming on soil-CO₂-release:

It is unquestioned that respiratory metabolism, including soil respiration, responds to temperature on an instantaneous, hourly or daily scale.

Hypothesis:

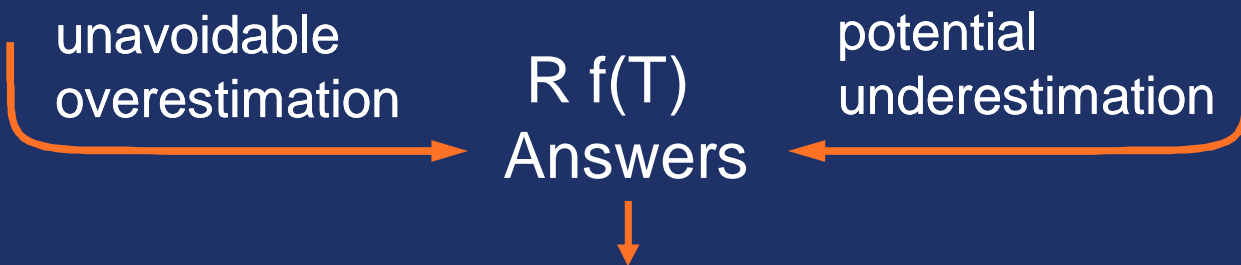
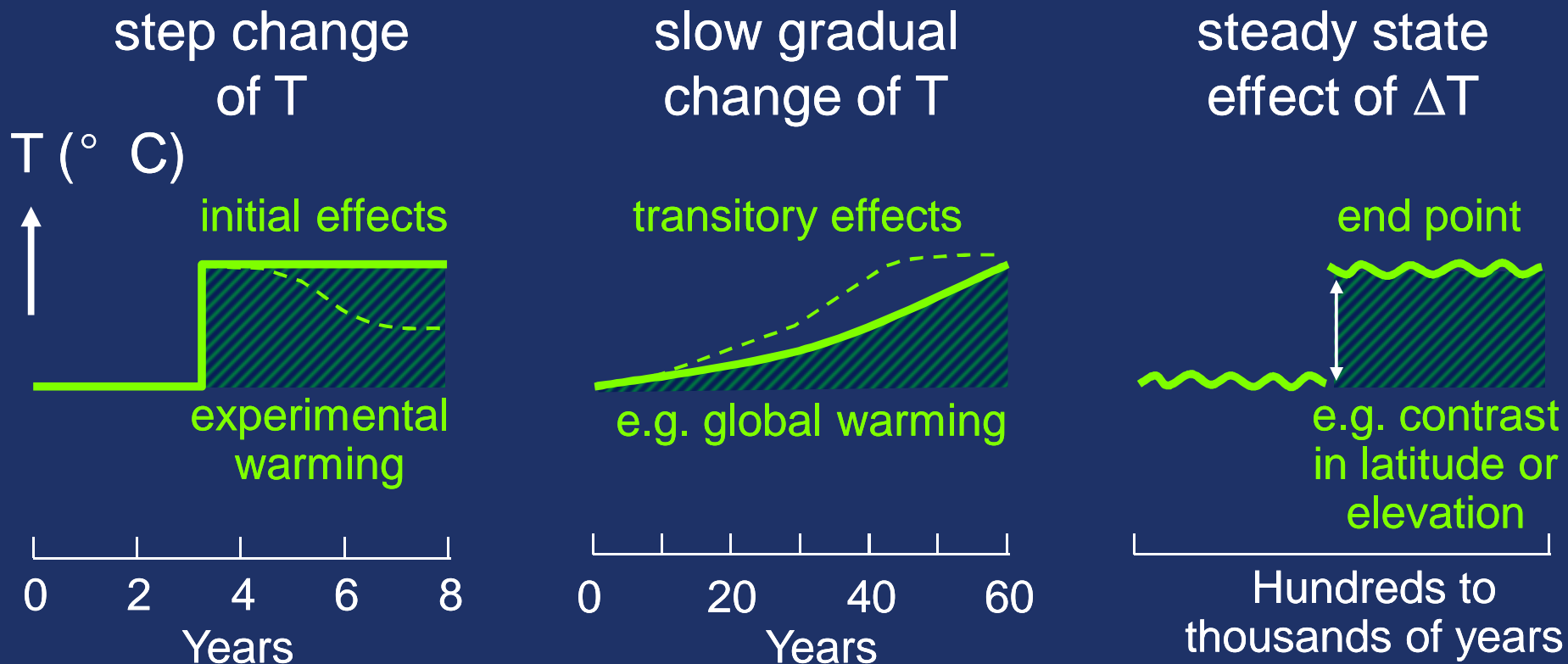
On a full or multi-year scale, soil respiration is a function of NPP, irrespective of temperature.

(Raich JW & Nadelhoffer KJ 1989, Ecology 70:1346
Kuzyakov Y & Gavrichkova O 2010, GCB 16:3386)

In summary,
I suggest a shift in paradigms

sink activity (demand) drives A
NPP (supply) drives R

Manipulative experiments



How many years does it take to arrive at steady state R?
Results sofar suggest $> 5 \ll 100$ years

The challenge:

How can we separate the intrinsic effect of rising temperature on R_s from the NPP-only driven effect on R_s ?

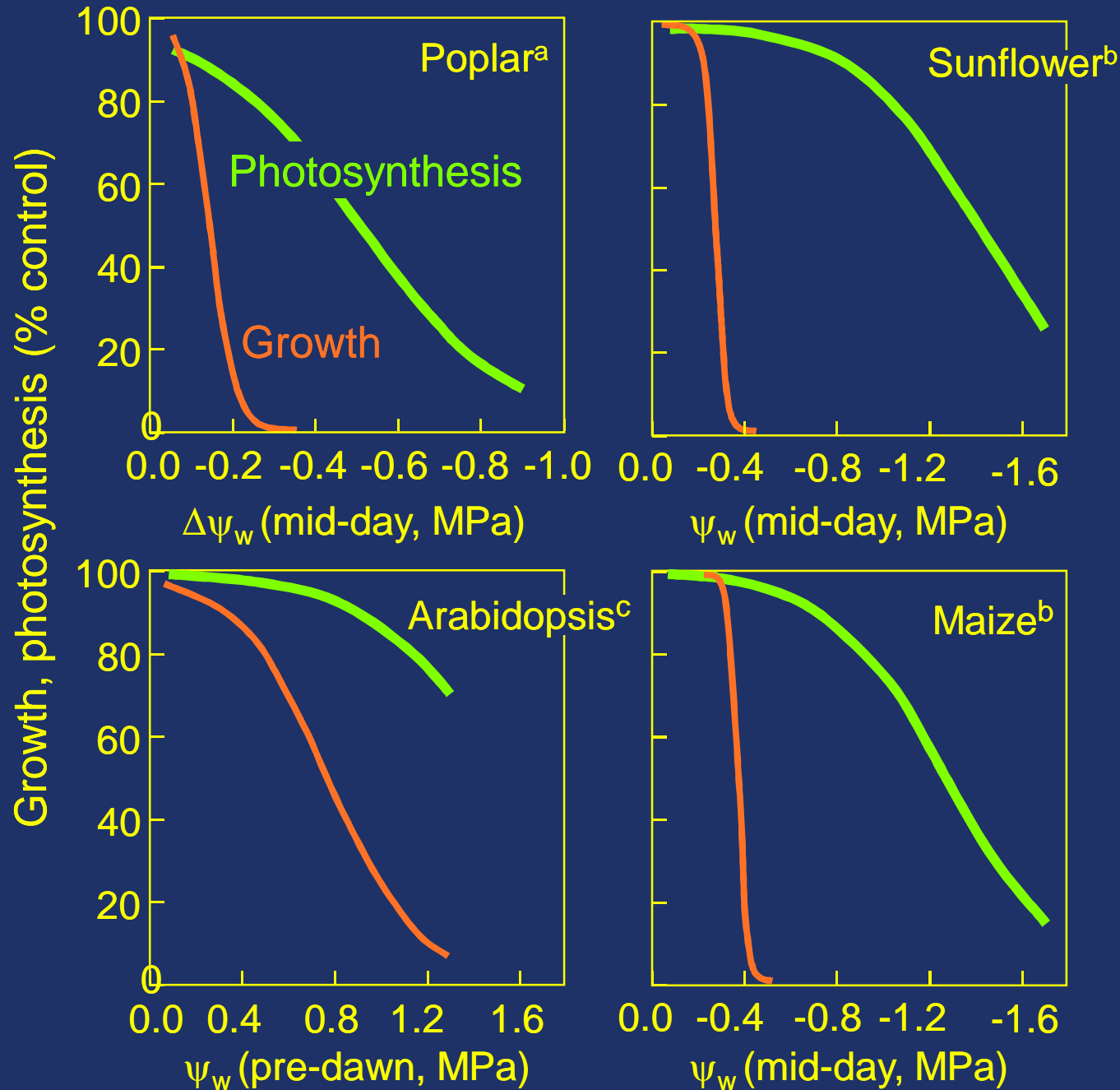
- ／ A warmer world may increase R_s , irrespective of NPP (net loss of C)
- ／ A warmer world may increase NPP (C-input) which in turn would stimulate R_s (with no additional T-effect on R_s)

- The only way soil respiration can be de-coupled from NPP in the long run, is a break-down of recalcitrant organic carbon in soils, carbon that is hundreds or thousands of years old.
- Since that break-down - if it does occur - would be very slow, the $R_s = f(\text{NPP})$ relationship would remain largely intact on a year to year basis even in a warmer future.
- on an hour to hour or day to day basis,
 $R = f(T, \text{H}_2\text{O})$

Both, drought and low temperature
affect sinks first, source activity follows



Differential sensitivity of shoot growth and photosynthesis to soil water deficit.



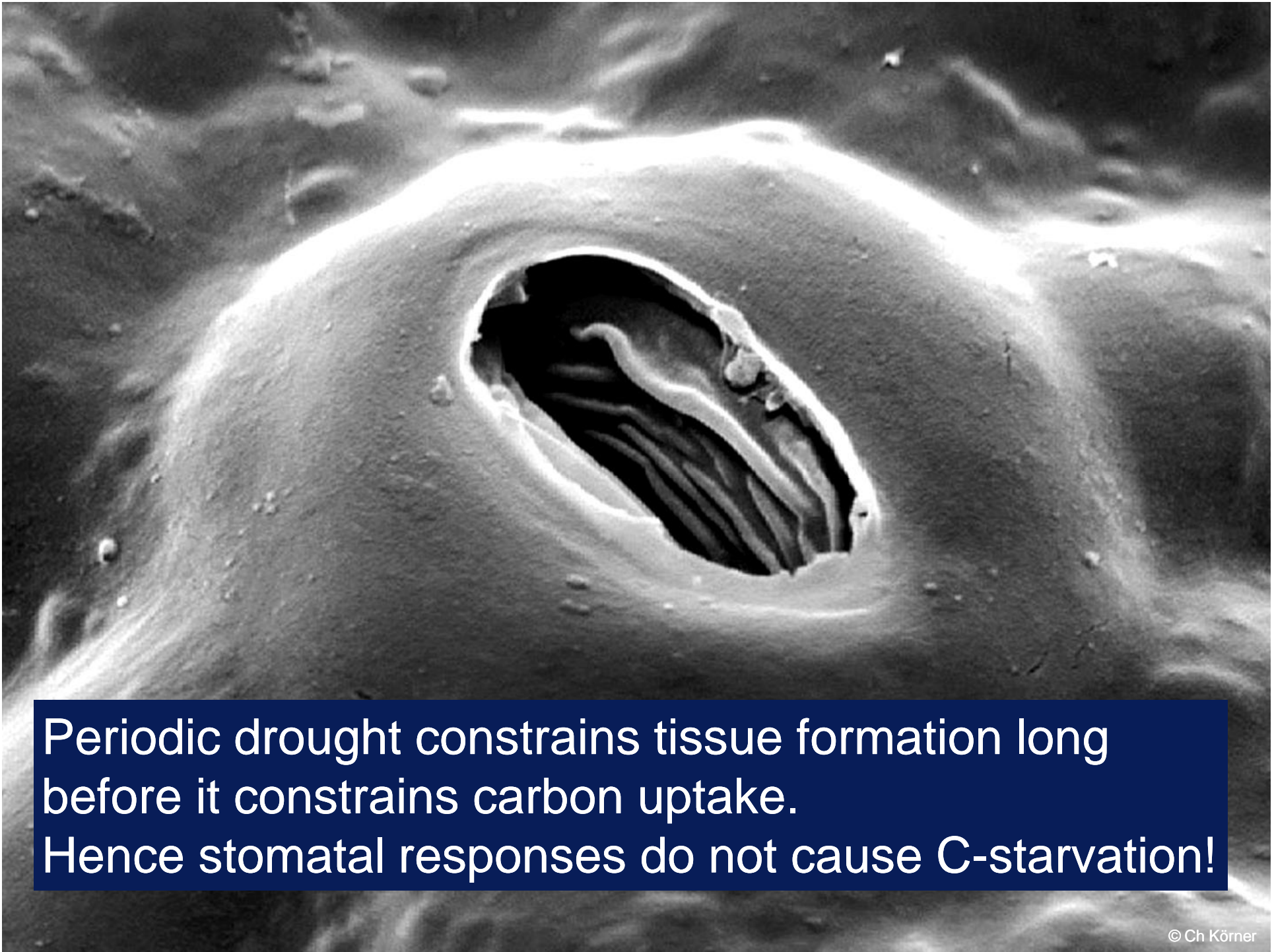
Muller B et al (2011)
J Exp Bot 62:1715

Data from :

^aBogeat-Tribouillot *et al* (2007)

^bBoyer (1970),
Tardieu *et al* (1999)

^cGranier *et al* (2006),
Hummel *et al* (2010)



Periodic drought constrains tissue formation long before it constrains carbon uptake.
Hence stomatal responses do not cause C-starvation!

Wet season

Dry season

Seasonal variation of NSC in tropical tree species

sun shade

5.8 ± 0.7 5.1 ± 0.8

6.0 ± 0.9 4.7 ± 0.7

shade sun

7.7 ± 0.8 $p = 0.002$ 7.2 ± 0.9 $p = 0.048$

8.4 ± 0.7 $p = 0.001$ 8.8 ± 1.1 $p = 0.094$

Leaves

Branches

Sink limitation:

Drought affects tissue formation first, photosynthesis last.

8.5 ± 1.8

9.7 ± 1.5 $p = 0.25$

Stem

→ NSC ↑

6.4 ± 1.2

8.3 ± 1.3

Coarse roots

2.2 ± 0.4

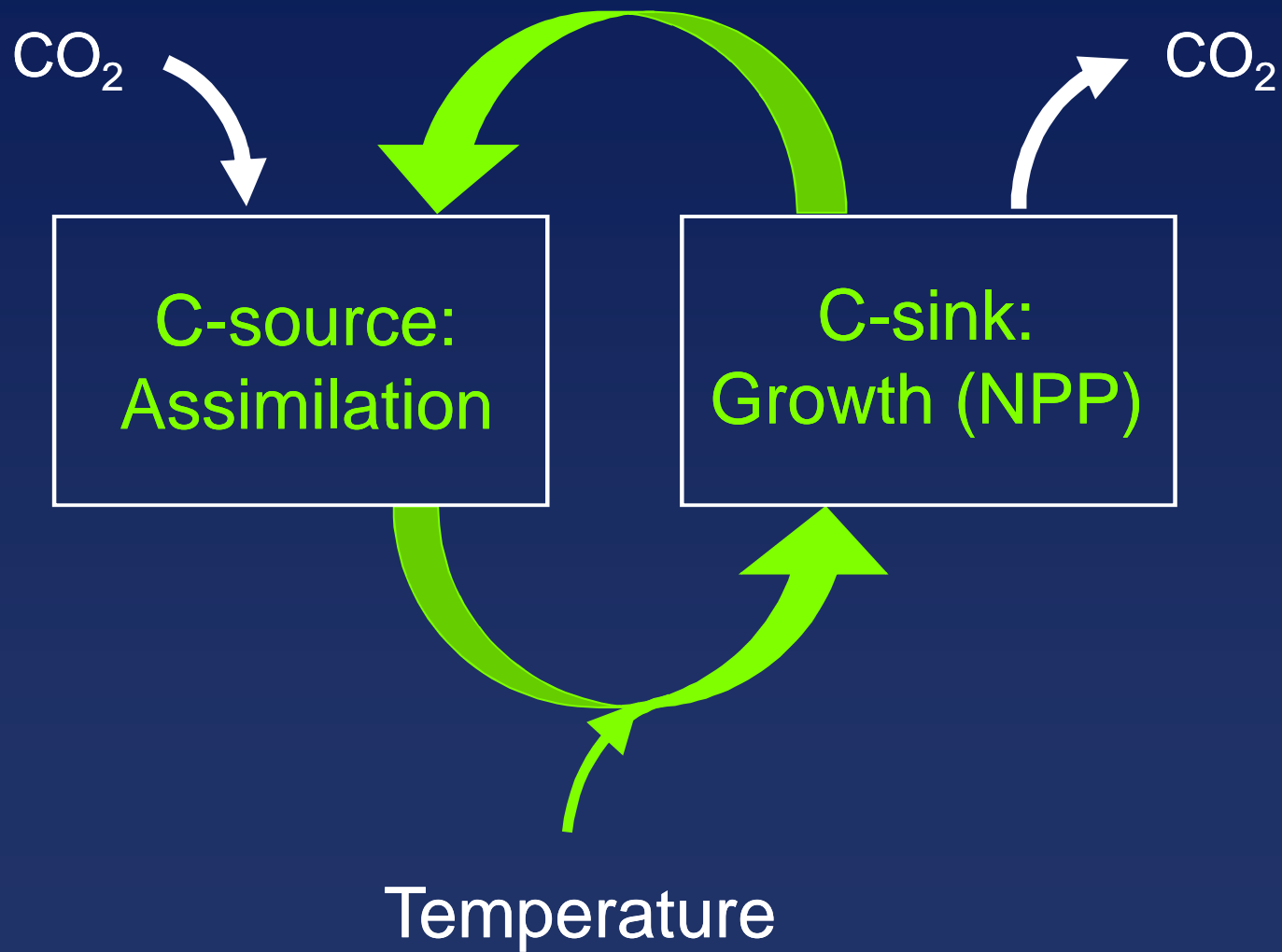
3.9 ± 0.8 $p = 0.012$

Fine roots

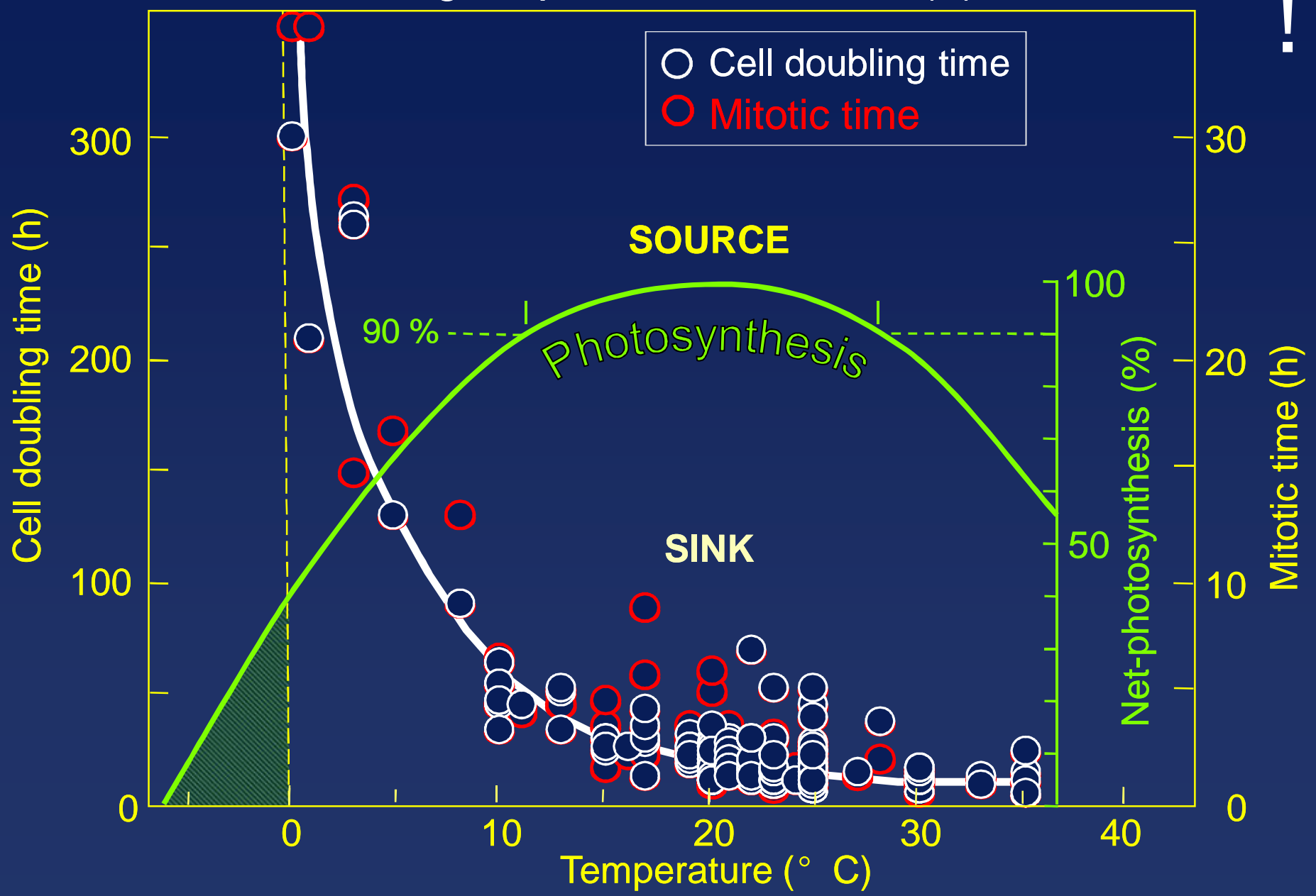
M Würth *et al* (2005)
Oecologia 143:11

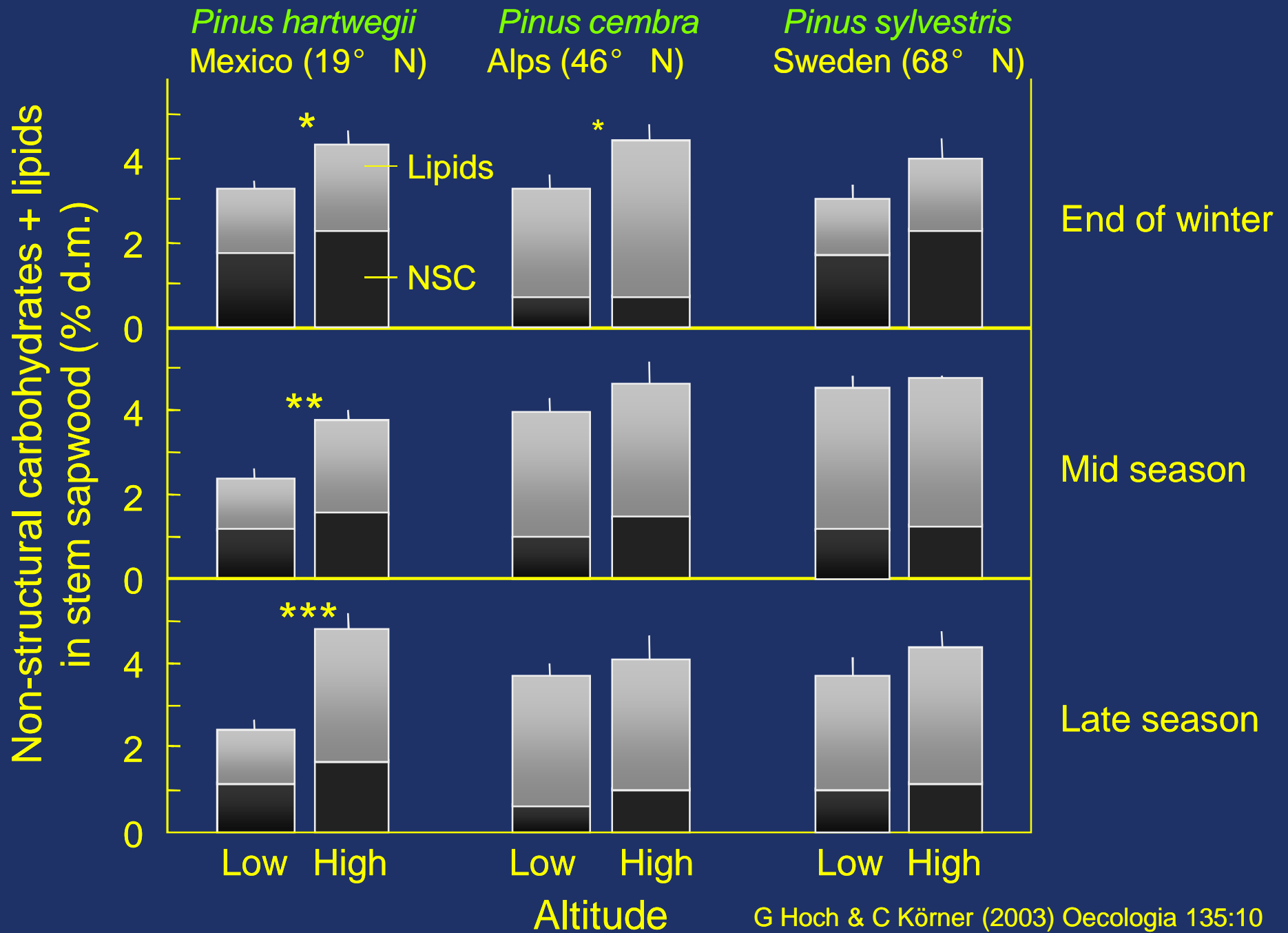


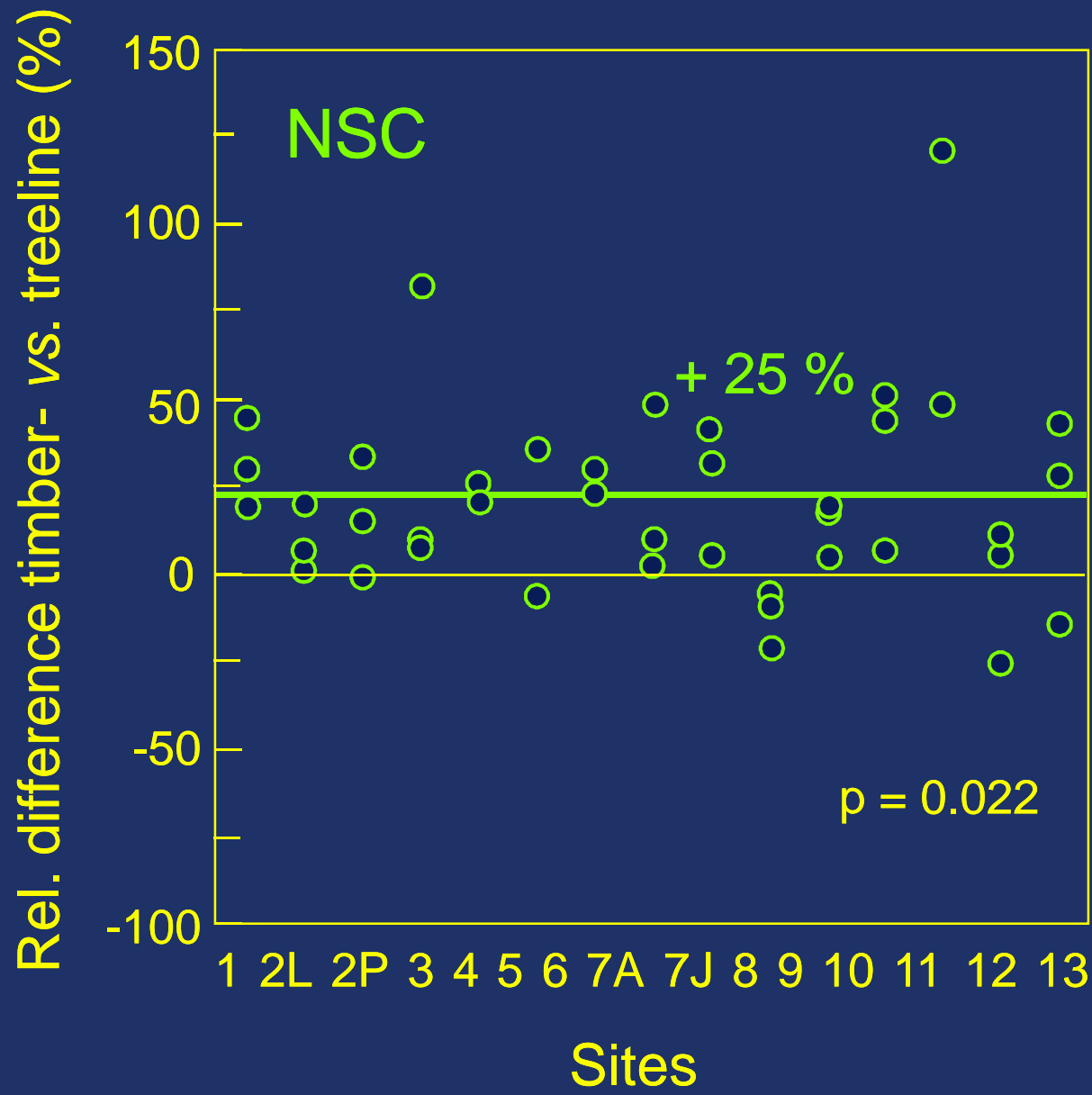
Low temperature limitation at the high elevation treeline:
Carbon shortage or constraints to carbon investment?



Higher plant cell division f (T)





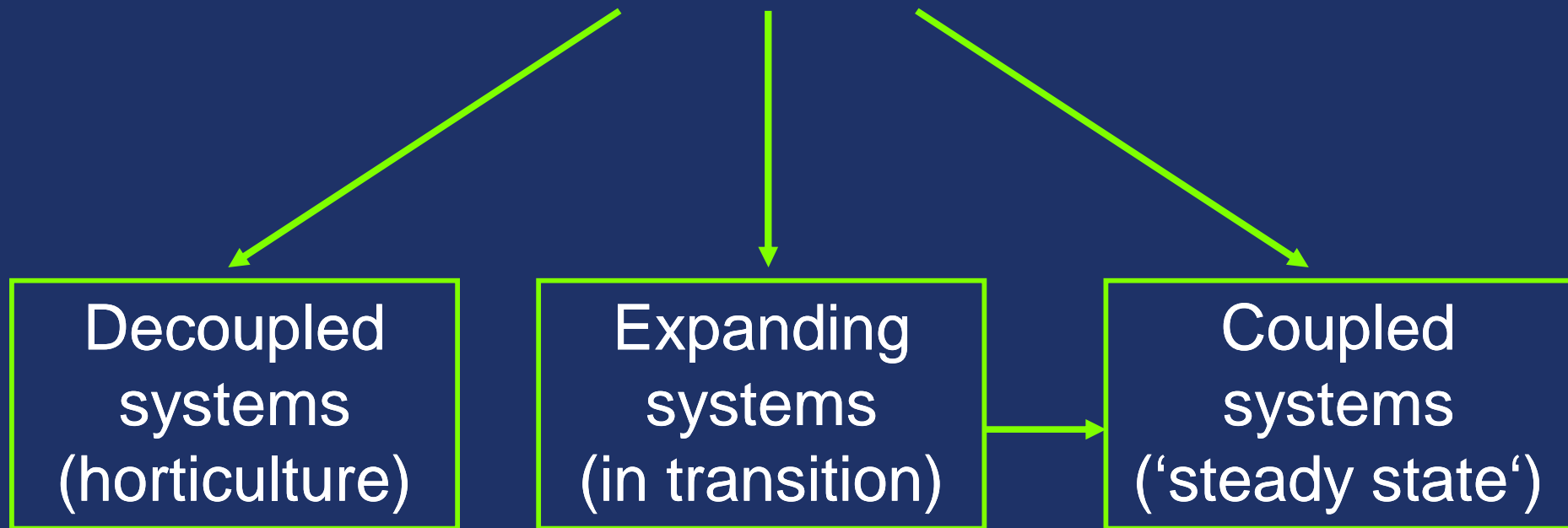


Classics in agrosience:

The 'magic' 6 ° C threshold

- "There seems to be a 6 ° C threshold temperature for plant development" (De Candolle, 1855).
- "The advent of spring may properly be considered as taking place at the advent of a 6-7 ° C isotherm" (Harrington, 1894).
- "6 ° C is considered the zero point of vital temperature" (USDA, Smith, 1920).

Plant growth in elevated CO₂



Coupling: - steady state fine root turnover
- steady state N-cycle
- steady state litter production (LAI)

Ch Körner (2006) Tansley Review, New Phytol 172:393



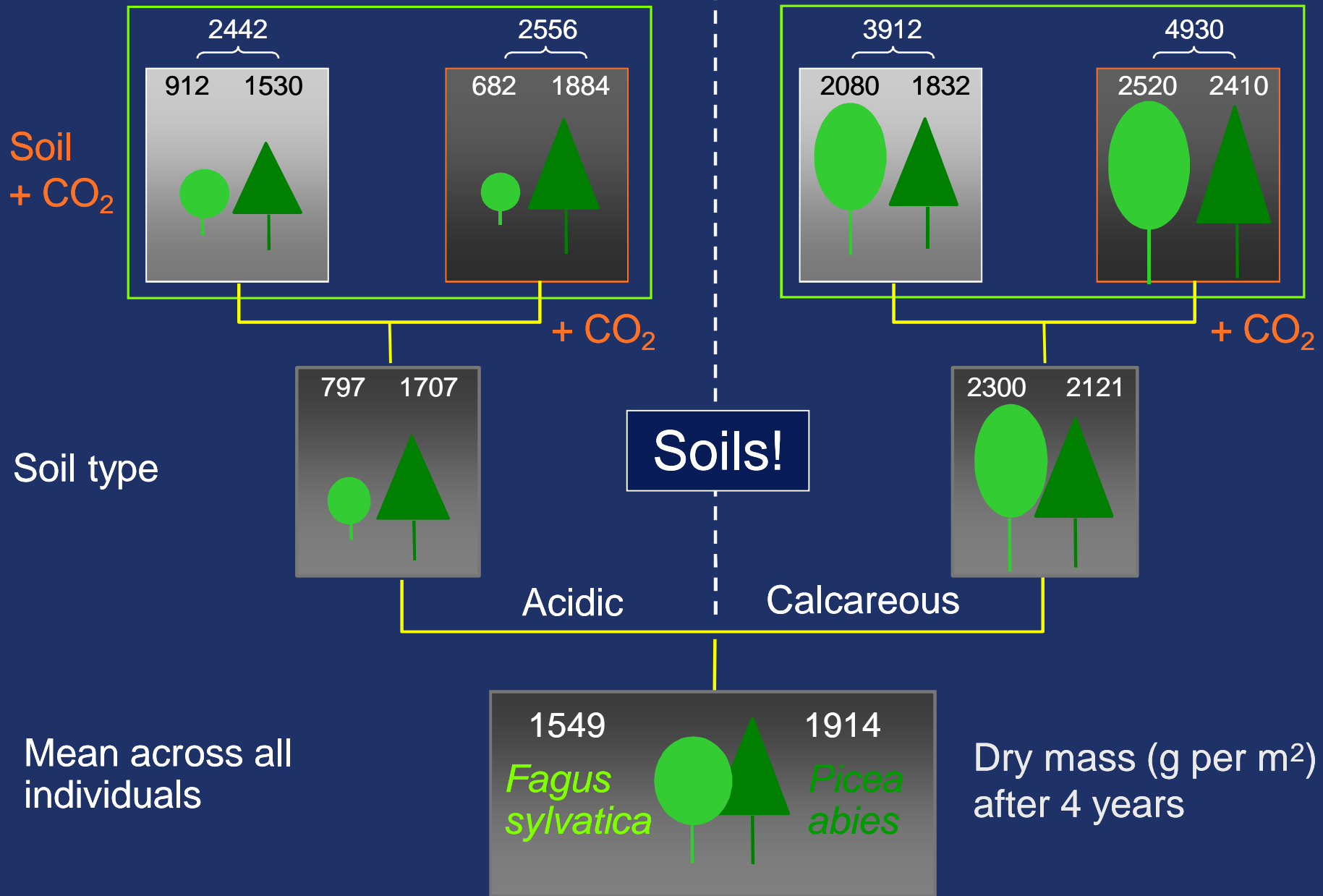
Decoupled horticultural system



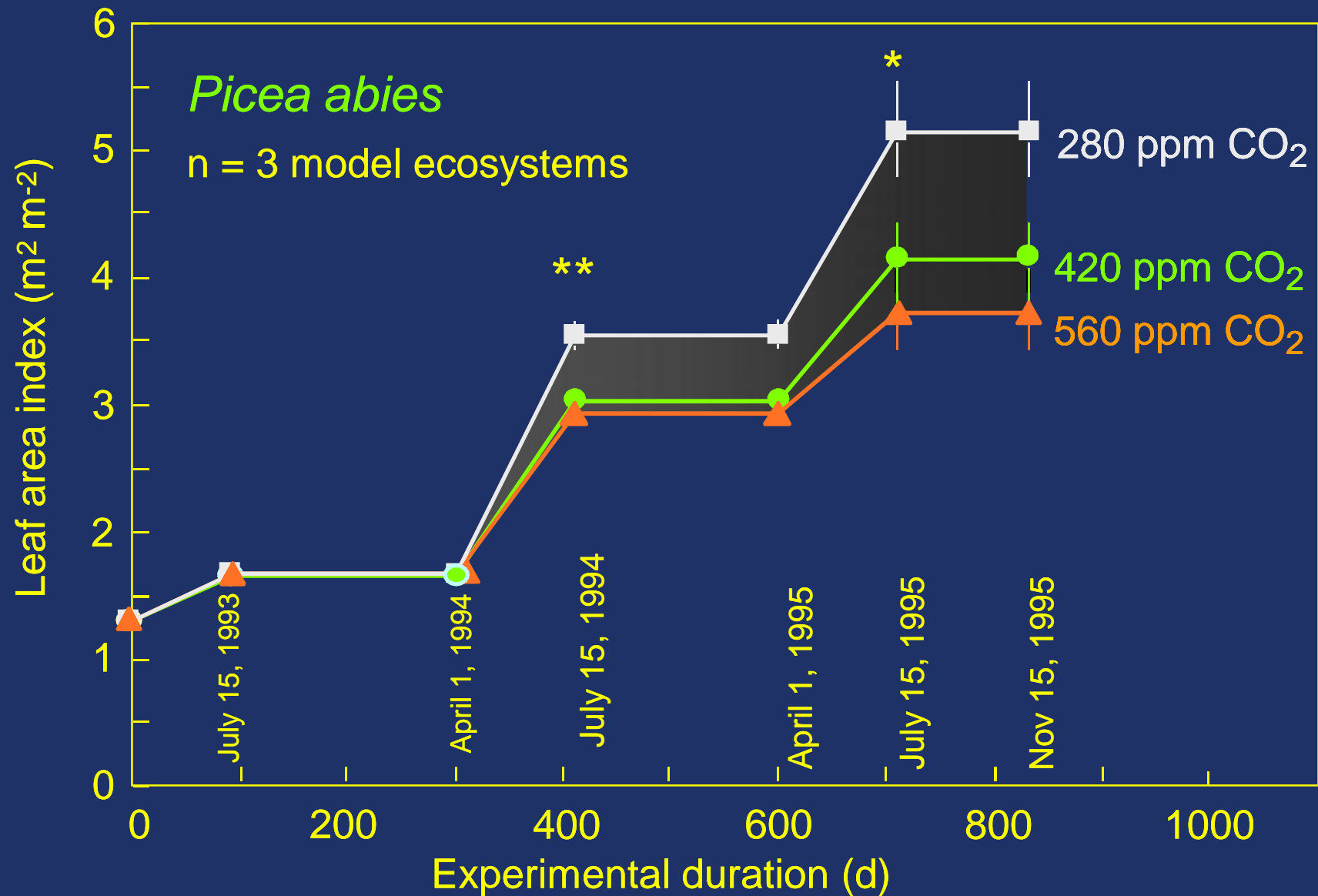
Expanding system

Fagus ↓

Fagus ↑



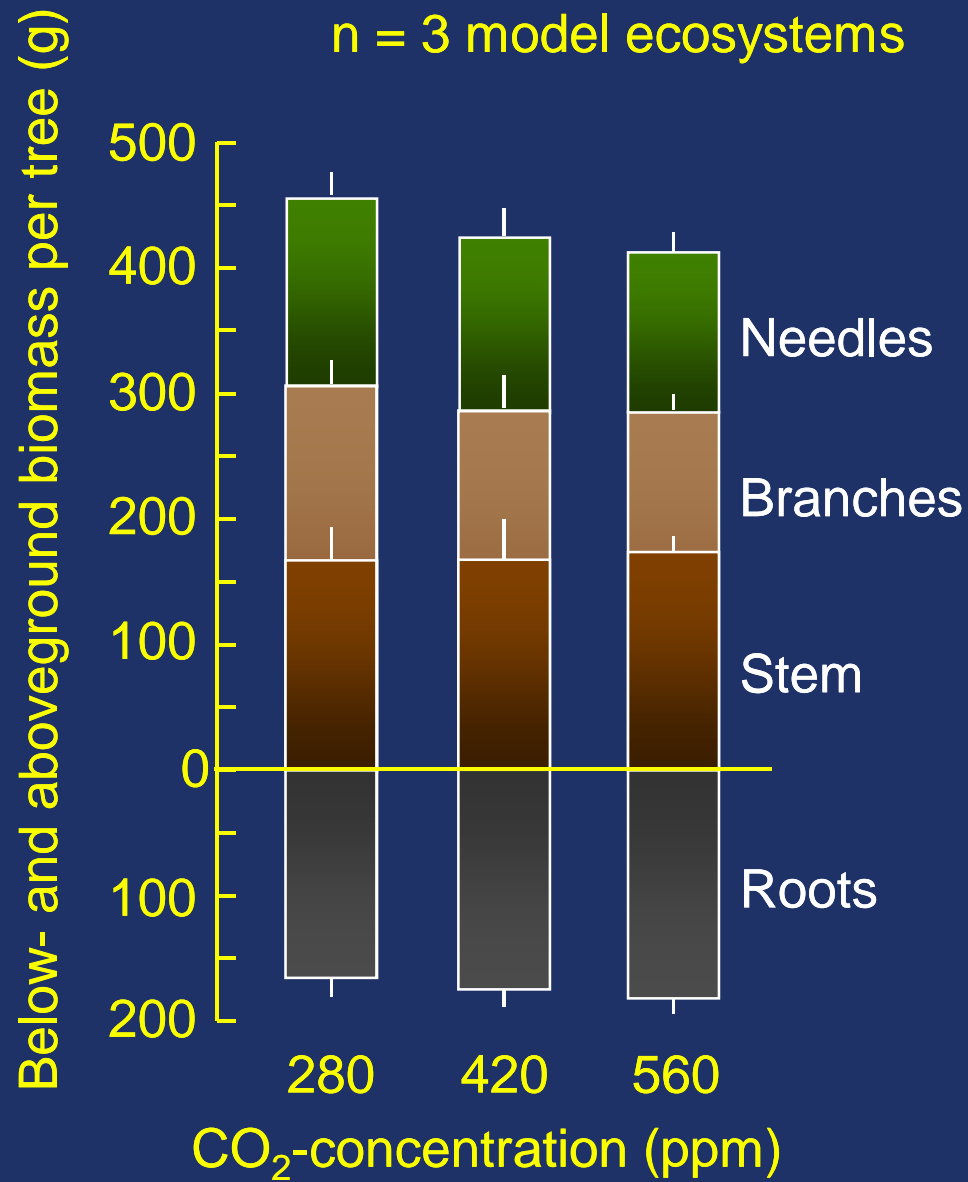
CO₂-induced nutrient shortage: LAI ↓



S Hättenschwiler & Ch Körner (1998) Oecologia 113:104

Picea abies biomass

n = 3 model ecosystems





Pinus taeda

Contradictory responses
to elevated CO₂ after
>10 years

Duke FACE:
stem radial growth
enhanced



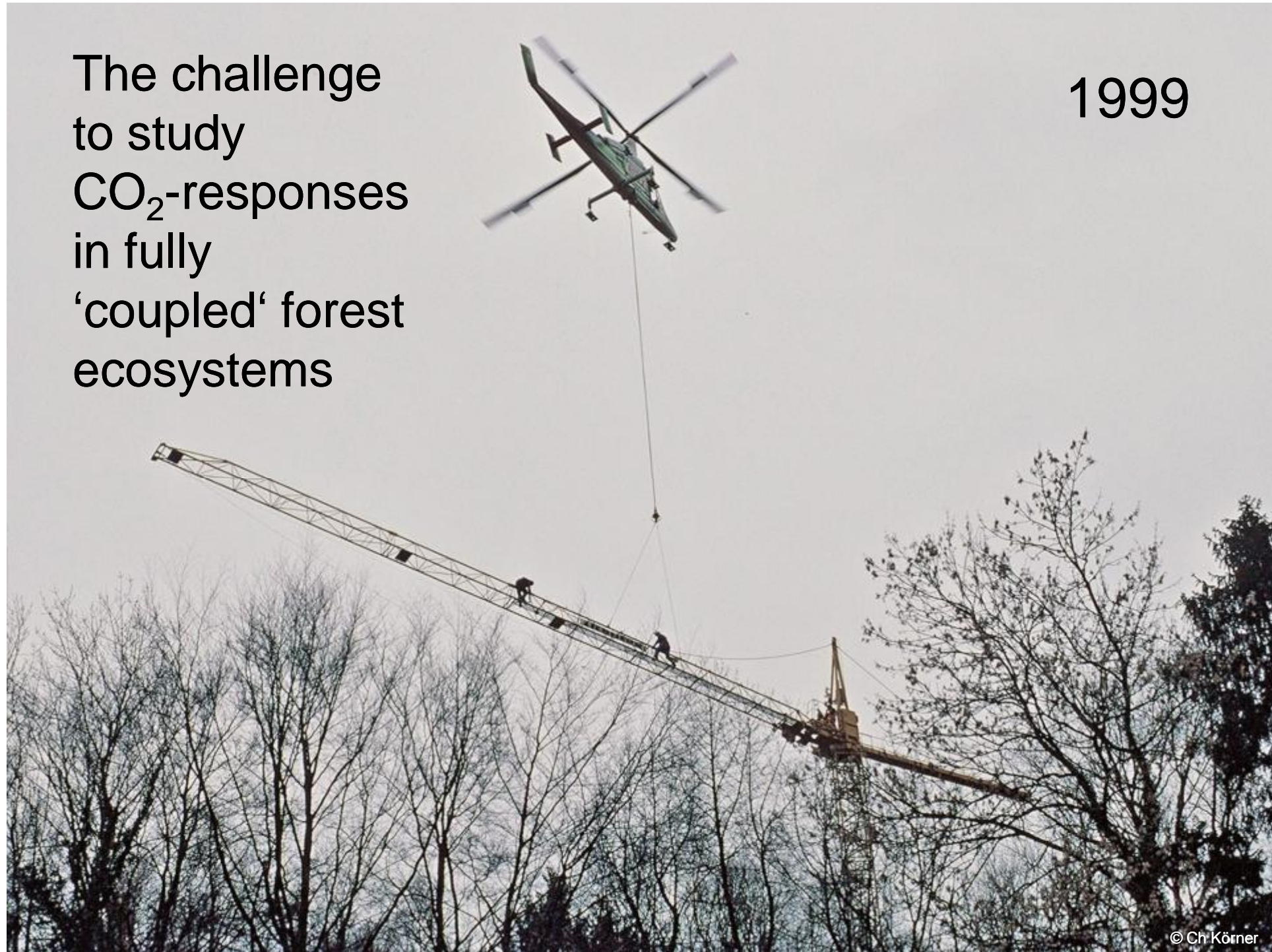
Liquidambar styraciflua

Oak Ridge FACE:
no effect
(roots responding in the
early phase)

Species or site effect?

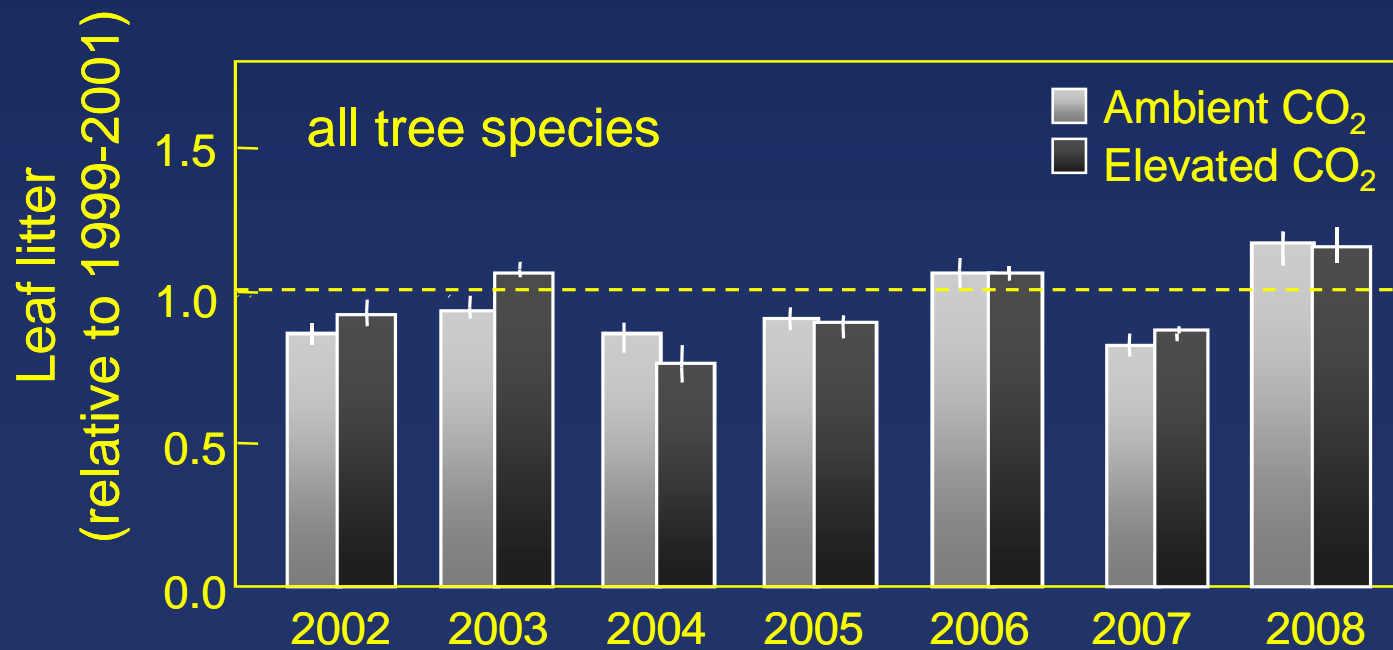
The challenge
to study
CO₂-responses
in fully
'coupled' forest
ecosystems

1999





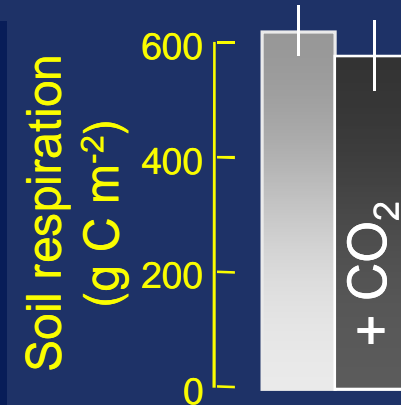
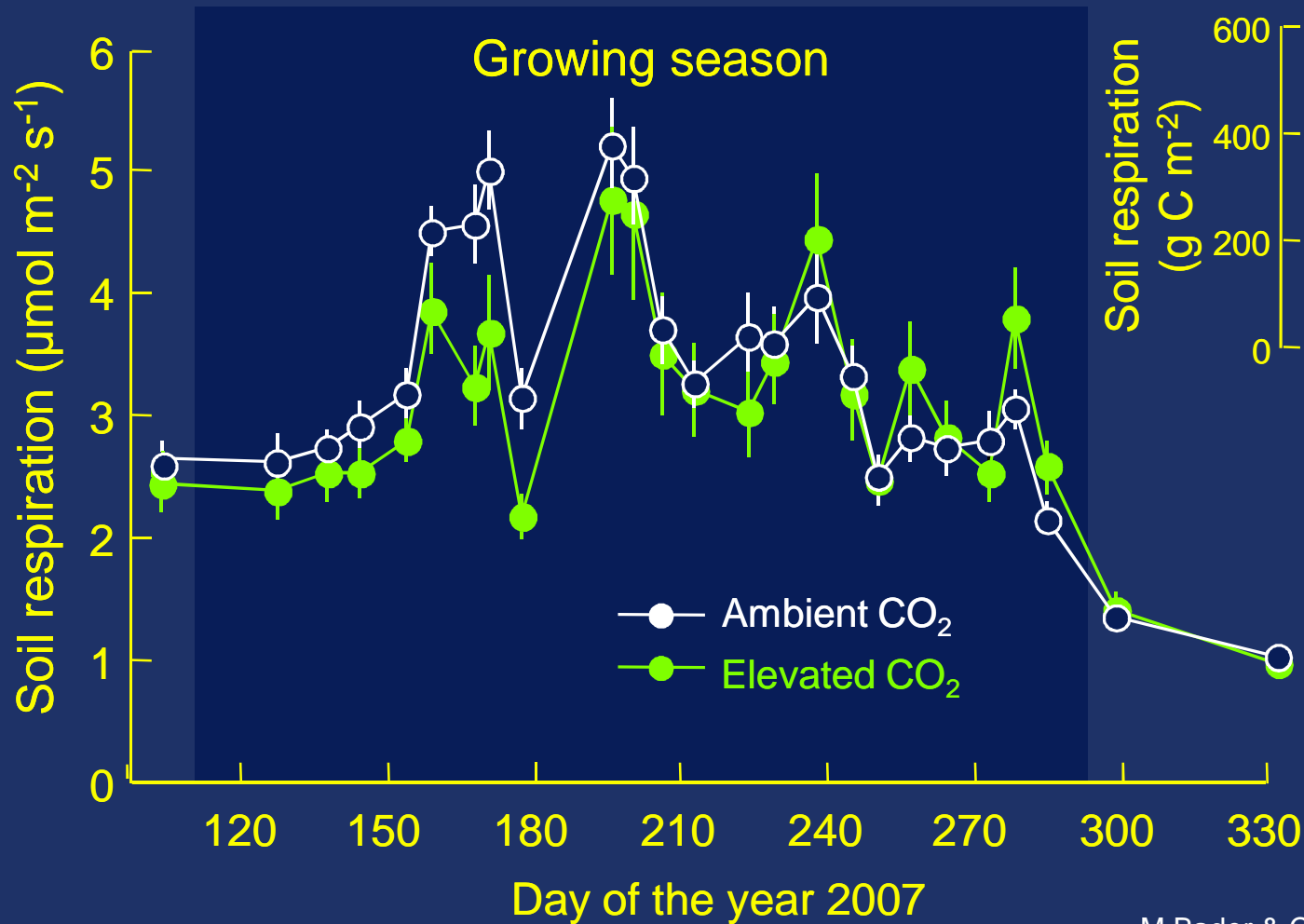
Leaf litter



No response

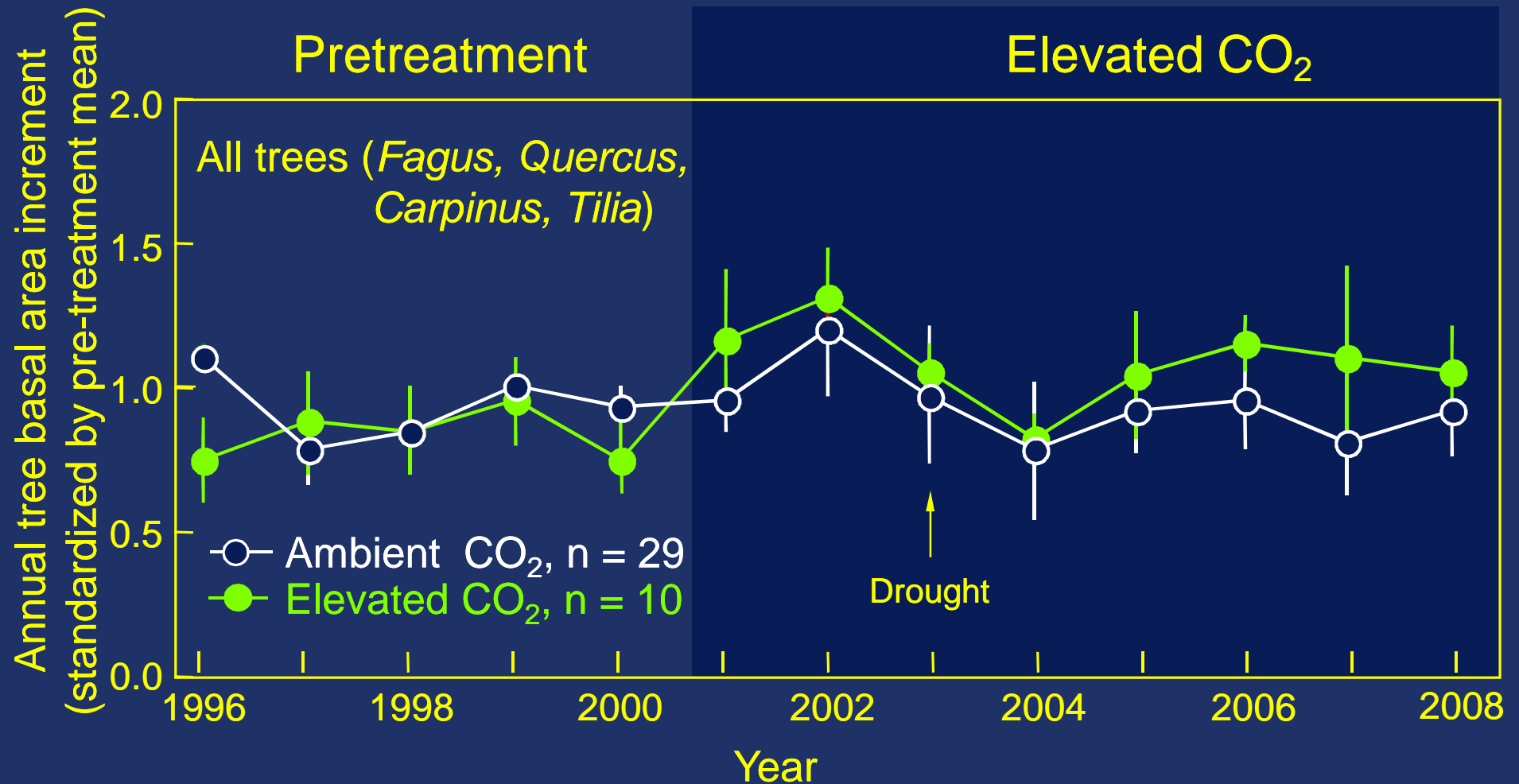
Ch Körner *et al.* (2005)
Science 309:1360 and
new data

CO₂ - release from soil - no response after 7 years



M Bader & Ch Körner (2010)
Glob Change Biol 16:2830

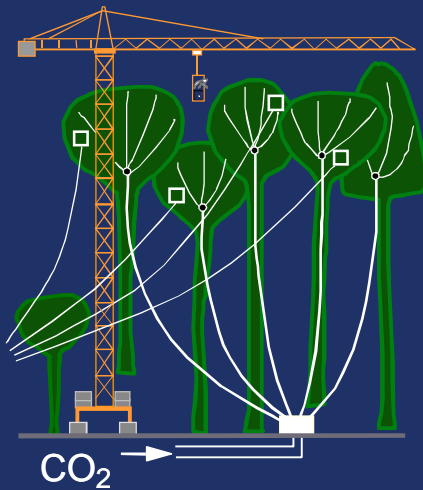
Growth of 100 year old trees in elevated CO₂, Swiss web-FACE



Ch Körner *et al.* (2005) *Science* 309:1360 and new data

Stomatal response to elevated CO₂

Swiss web-FACE



- None



Quercus

- Weak

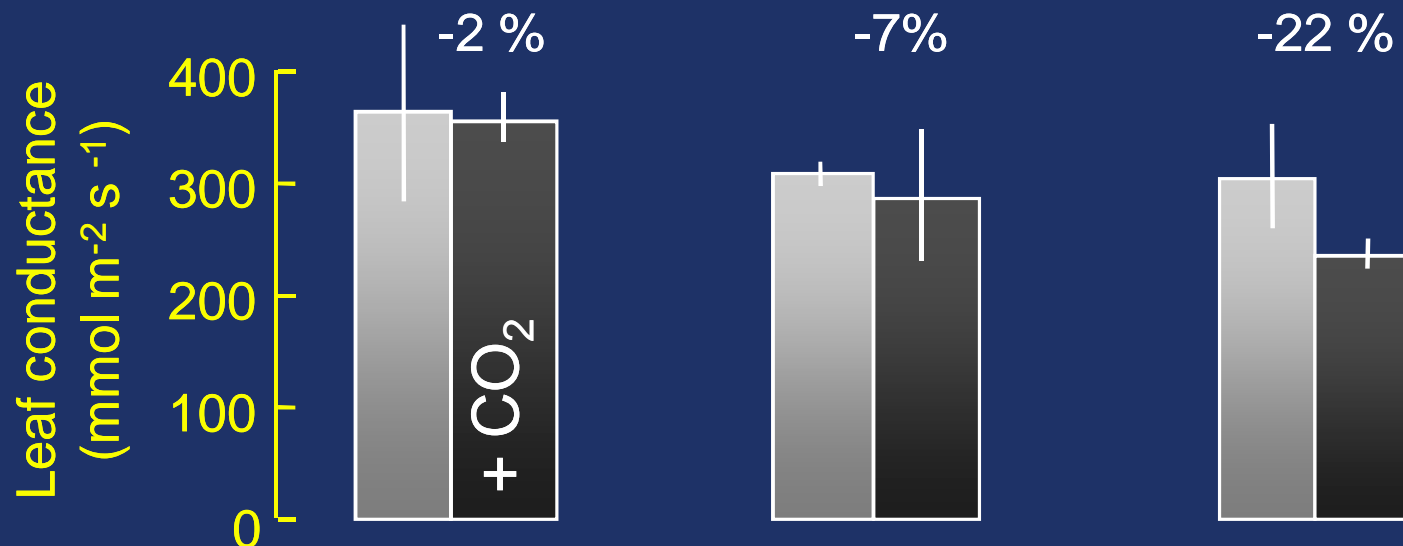


Fagus

- Strong



Carpinus



S Keel *et al* (2007) *Trees* 21:151

Species differ in their responses



Biodiversity effects
(winners and losers, new assemblages)

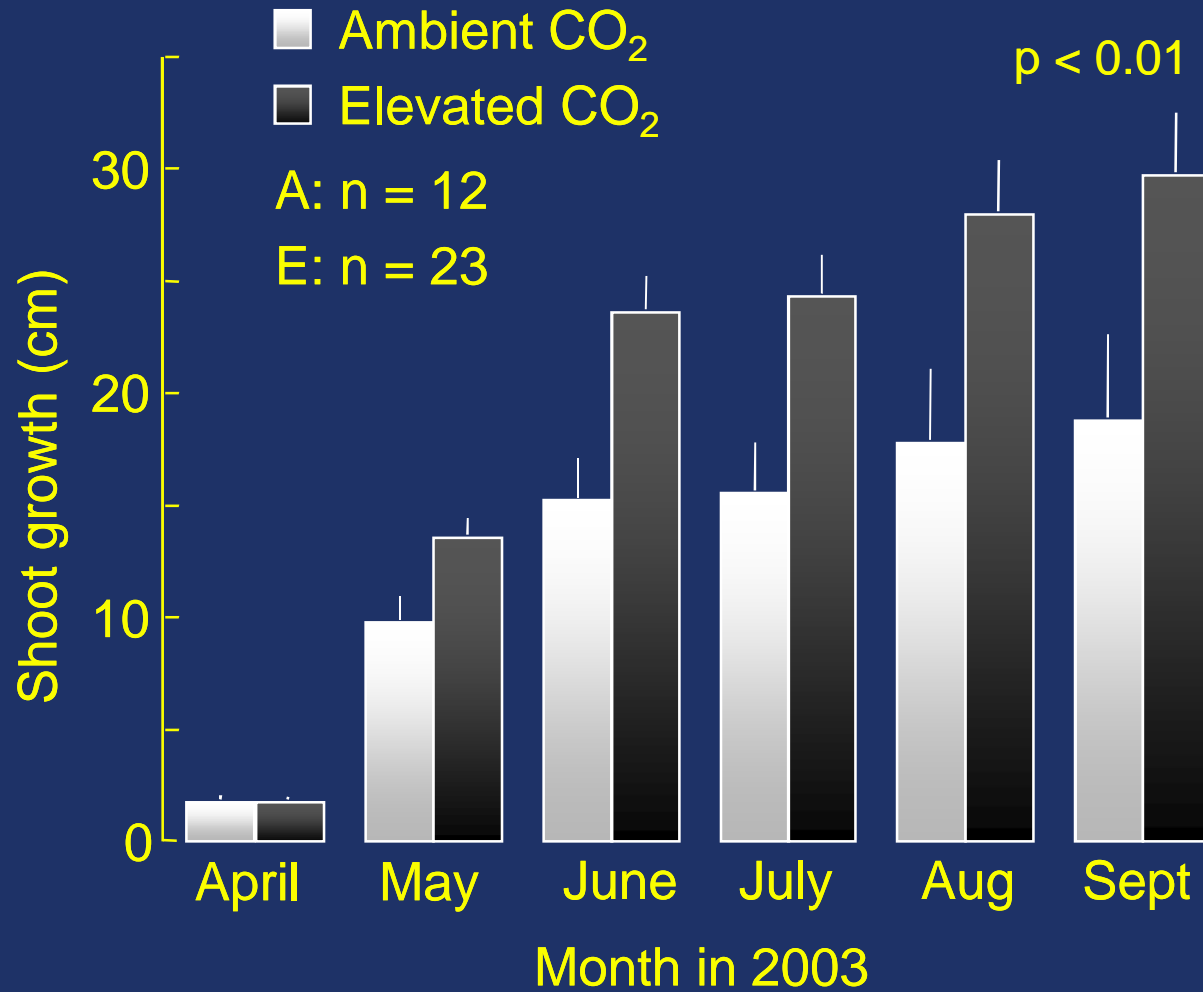


Ecosystem effects
(water loss, carbon and nutrient cycle)

→ 2 examples

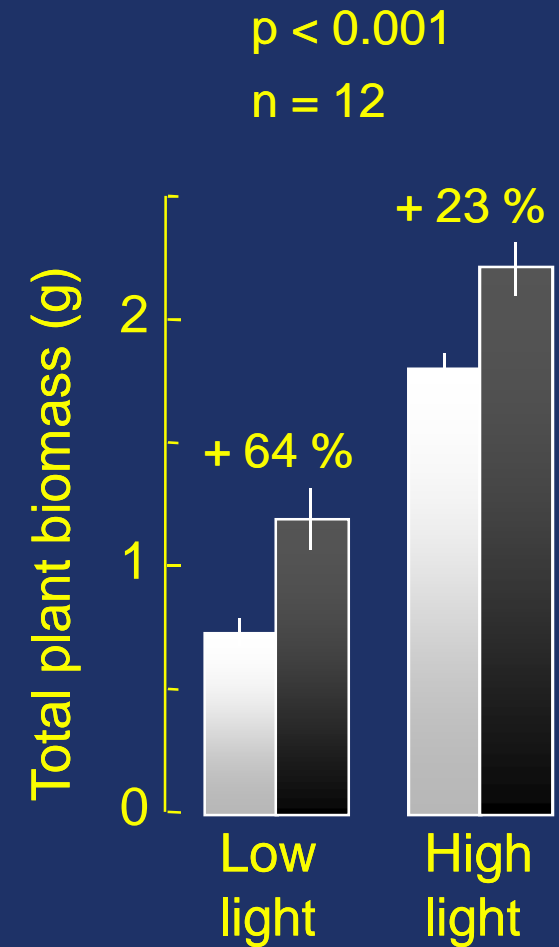
Hedera helix

Understorey, Swiss web-FACE



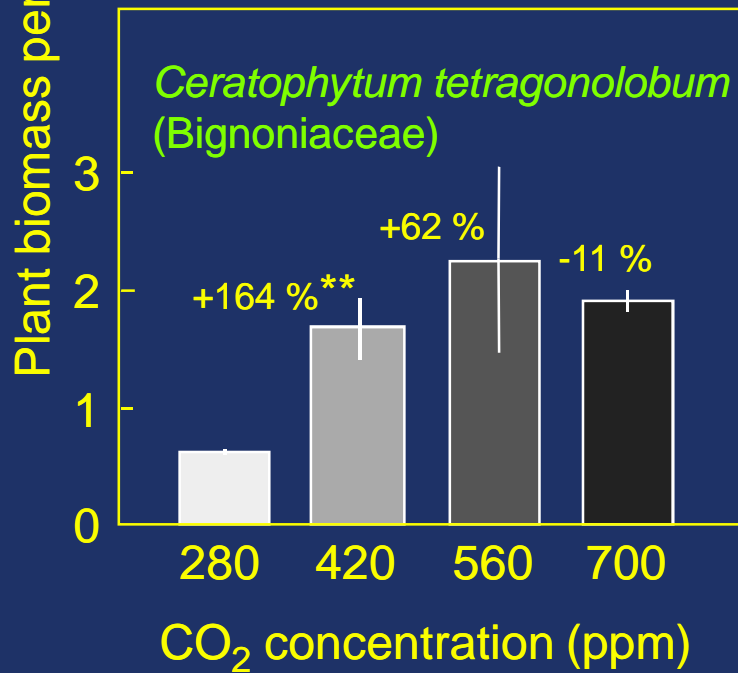
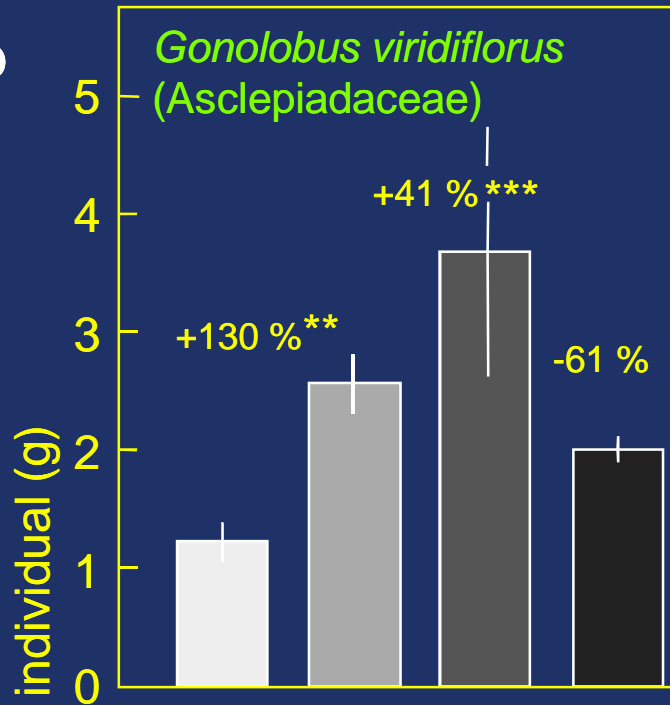
G Zotz *et al* (2006) *Funct Ecol* 20:763
Ch Körner (2006) *Biologist* 53:2

Clematis vitalba



F Grob (unpubl.)

In deep shade, lianas are gaining



J Granados & Ch Körner (2002) GCB 8:1109
Ch Körner (2003) Phil Trans R Soc Lond A 361:2023

Summary CO₂ effects

- **First principle responses** (e.g. photosynthetic or stomatal CO₂-responses) might become **overrun by biodiversity effects** in the long term.
- Elevated atmospheric CO₂ is unlikely to cause a sustained increase of **carbon storage** in biomass even if it stimulates growth.
- All experimental data are confounded by stomata mediated water effects.
- Deep shade is the most likely place to see positive CO₂-signals.

Overall summary

- Carbon capture is controlled by C-sinks.
- C-sinks are controlled by soil resources (nutrients, water, stoichiometry constraints).
- Ecosystem respiration is controlled by substrate availability (NPP) rather than T.
- Drought and low temperature act upon sink activity.
- Ecosystem CO₂-saturation because sink activity controls source activity and because soil resources are finite.
- Life in deep shade is intrinsically C-limited.
- Biodiversity and its change affects the C-cycle.